

Article

Travel-Time Inversion Method of Converted Shear Waves Using RayInvr Algorithm

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Abstract: The detailed studies of converted S-waves recorded on the Ocean Bottom Seismometer (OBS) can provide evidence for constraining lithology and geophysical properties. However, the research of converted S-waves remains a weakness, especially the S-waves' inversion. In this study, we applied a travel-time inversion method of converted S-waves to obtain the crustal S-wave velocity along the profile NS5. The velocities of the crust are determined by the following four aspects: (1) modelling the P-wave velocity, (2) constrained sediments V_p/V_s ratios and S-wave velocity using PPS phases, (3) the correction of PSS phases' travel-time, and (4) appropriate parameters and initial model are selected for inversion. Our results show that the v_s and V_p/V_s of the crust are 3.0–4.4 km/s and 1.71–1.80, respectively. The inversion model has a similar trend in velocity and V_p/V_s ratios with the forward model, due to a small difference with ΔV_s of 0.1 km/s and $\Delta V_p/V_s$ of 0.03 between two models. In addition, the high-resolution inversion model has revealed many details of the crustal structures, including magma conduits, which further supports our method as feasible.

Keywords: converted S-waves; travel-time inversion; RayInvr; ocean bottom seismometer (OBS)



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1. Introduction

The Ocean Bottom Seismometer (OBS) survey is a kind of seismic observation technology whose geophones are directly deployed on the seafloor, which has a wide range of use in the research of deep seismic structures [1,2]. In contrast to Multichannel Seismic Reflection (MCS) survey, the OBS survey has wider detection coverage, both in depth (>25 km), and width (>100 km) [3]. Another advantage is the ability to record converted S-waves [4,5], because OBSs, including 4-component geophones, are placed on the seafloor. The joint studies of P- and S-wave velocities derived V_p/V_s ratios of the crustal structures, which provide direct evidence for constraining the lithology and geophysical properties [6,7]. Therefore, the S-wave research has made a huge contribution to distinguishing crystalline basement compositions [4,8–10] and mantle serpentinization [11], as well as estimating the saturation of gas hydrates or fluids [12]. However, the RayInvr forward approach, one of the common modelling methods of the converted shear wave, has complicated steps and leads to difficulties in ensuring the optimal solution, which is urgently needed to improve the existing methods.

The detailed studies of converted S-waves remain a weakness. One reason for this is that the converted S-waves are hard to observe due to the complicated submarine environment (the coupling between the instruments and seabed) and the specific conversion interface for P- to S-wave. Previous studies have mainly been limited to modelling P-wave velocity structure, and only three OBS profiles (Figure 1) have revealed the S-wave velocity

structure in the north-eastern continental margin of the South China Sea (SCS) [4,13,14]. Additionally, the research of converted S-waves has a critical defect because of the deficient modeling method. For example, only the V_p/V_s ratio of the entire layer in the model was obtained by using 2-D ray tracing along the profile on the Lofoten continental margin, northern Norway [15]. The same cases were seen by Chian and Loudon in 1994 [8], and Kodaira et al. in 1996 [16], too. Almost all similar studies used the RayInvr forward modelling [17] after 2003 [9], when the forward approach was first applied to modelling the observed converted S-waves data [4,10,11,18]. In the RayInvr model, the depth and velocity nodes of layers defined variable-sized trapezoidal blocks. In that case, only Poisson's ratio (σ) of different blocks was adjusted to acquire the forward S-wave model. The forward modelling method is often affected by human subjectivity (e.g., if the grid size of blocks is set too small, it will result in a huge workload; otherwise the model resolution will be insufficient), so a more objective inversion method is needed.

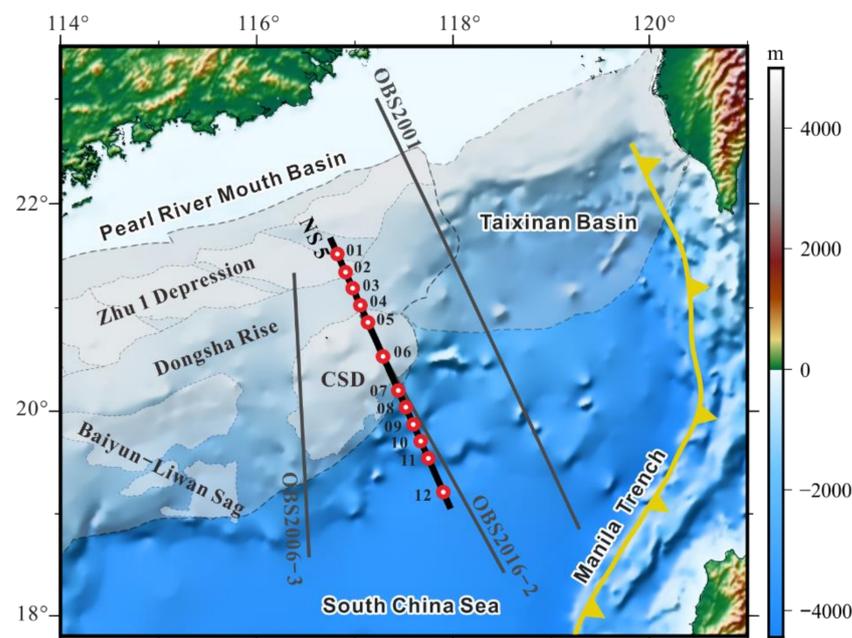


Figure 1. Topography and bathymetry map of north-eastern South China Sea (SCS) with the location of deep seismic survey lines. The gray lines represent the deep seismic survey lines which have conducted the studies of converted S-waves, including OBS2001 [4], OBS2006-3 [13], OBS2016-2 [14]. The thick, black, solid line is the profile NS5 in this study. Red dots represent ocean bottom seismometers (OBSs). Gray areas represent the distribution of the Pearl River Mouth Basin and Taixinan Basin. The yellow curve with triangles denotes Manila Trench. CSD: Chaoshan Depression.

In May–July 2014, a deep seismic refraction profile (NS5) was carried out in the north-eastern SCS (Figure 1). A large number of converted S-wave phases through the crust were recorded by the OBSs, which provided good data for crustal S-wave velocity inversion. In this study, we introduced the improvements in the inversion method to acquire the crustal S-wave velocity model using converted S-waves phases.

2. Data and Methodology

2.1. Data Acquisition and Processing

A wide-angle seismic profile (NS5) was conducted by the R/V Shiyan 2 of the South China Sea Institute of Oceanology in May–July 2014. The profile NS5 is 317 km long and is an NW-SE trending line crossing the Dongsha Rise (Figure 1). OBSs are deployed at an interval of ~15–20 km along the profile with 12 OBS instruments (comprising three orthogonal geophones and one hydrophone) that record usable data. The seismic

source consisted of four BOLT air-guns arrayed with a total volume of 6000 in³ (~98.3 L), which were operated at a pressure of 2000 psi (13.79 MPa). A total of 965 shots were fired along profile with a shooting interval time of 120 s. The recorded OBS data are of high quality, and most instruments have recorded clear air-gun signals. The processing of OBS data includes data format conversion, correction of OBS position and clock drift, and thereby four components of seismic profile are plotted. We have identified five P-wave phases on the vertical component profile: refracted/reflected phases in the Cenozoic sediment layers (Ps2/Psp), refracted phases in the Mesozoic sediment layers (Ps3), refracted phases in the upper and lower crust (Pg), refracted phases in the upper mantle (Pn), and reflected phases at the crust-mantle boundary (PmP). Based on the seismic phases described above, we set seven layers in the initial velocity model (water, Cenozoic sediment, Mesozoic sediment, upper crust, lower crust, high-velocity lower crust, and upper mantle), corresponding to velocities of 1.5 km/s, 1.7–3.5 km/s, 3.5–5.5 km/s, 5.5–6.5 km/s, 6.5–7.0 km/s, 7.0–7.5 km/s, and 7.8–8.2 km/s. The Cenozoic sediment layer was further divided into two layers, each corresponding to a high-amplitude reflection in single-channel reflection seismic data [19] and OBS data (Figure 2). Furthermore, we performed forward P-wave velocity modelling using the RayInvr [17]. We adjusted the boundary nodes and velocity values by using a manual trial-and-error method to make the modelled, theoretical travel times consistent with observed times, based on ray tracing and travel-time fitting. The final model (Figure 3 and Figure S1) had the root-mean-square travel-time residuals (RMS) of 81 ms, and normalised chi-squared value (χ^2) of 1.468. The uncertainty of the Moho depth and P-wave velocity of the high-velocity lower crust were determined to be -0.2 – 0.2 km and -0.2 – 0.2 km/s (Figure S2), respectively. The converted S-waves phases are primarily concentrated in the two horizontal components. We rotated two horizontal components into radial and transverse components to make the S-wave phases clearer. The polarisation angle (for details see Table S1) is calculated using the energy scanning method [20]. Then, the S-wave modelling is conducted using the S-wave phases recorded on the radial component.

2.2. Method

In this study, we apply the RayInvr algorithm [17] for crustal S-wave velocity inversion. RayInvr allows the flexibility of ray tracing and inverts different converted S-waves phases, simulating different mode conversions. It also has the ability to switch between forward modelling and inversion, to allow V_p/V_s to vary within model layers.

Based on asymptotic ray theory of the zeroth order, the ray-tracing equations (Equation (1)) were solved by Runge–Kutta method [21] with error control, and Snell's Law is applied to model interfaces. In Equation (1), θ is the angle between the tangent of the ray and the z axis, v is the wave velocity, v_z and v_x are the partial derivatives of the velocity on the z and x axis, respectively. The initial values of x , z and θ are x_0 , z_0 (shooting coordinates) and θ_0 (take-off angles of rays).

$$\frac{dx}{dz} = \tan \theta \cdot \frac{d\theta}{dz} = \frac{v_z \tan \theta - v_x}{v} \quad (1)$$

Equation (2) is the discrete form of the travel-time equation, where l_i and v_i are the path length and velocity of the i -th ray segment. Equation (3) is derived from Equation (2) to linearize by Taylor series expansion, and the higher-order terms are neglected. In Equation (3), A is the partial derivative matrix, Δm is the model parameter adjustment vector, and Δt is the travel-time residual vector. The model parameter adjustment vector was calculated in each inversion and used to update the model. After multiple iterations, inversion will be terminated when the preset criteria are reached, then an optimum inverted model will be obtained.

$$t = \sum_{i=1}^n \frac{l_i}{v_i} \quad (2)$$

$$\Delta t = A\Delta m \tag{3}$$

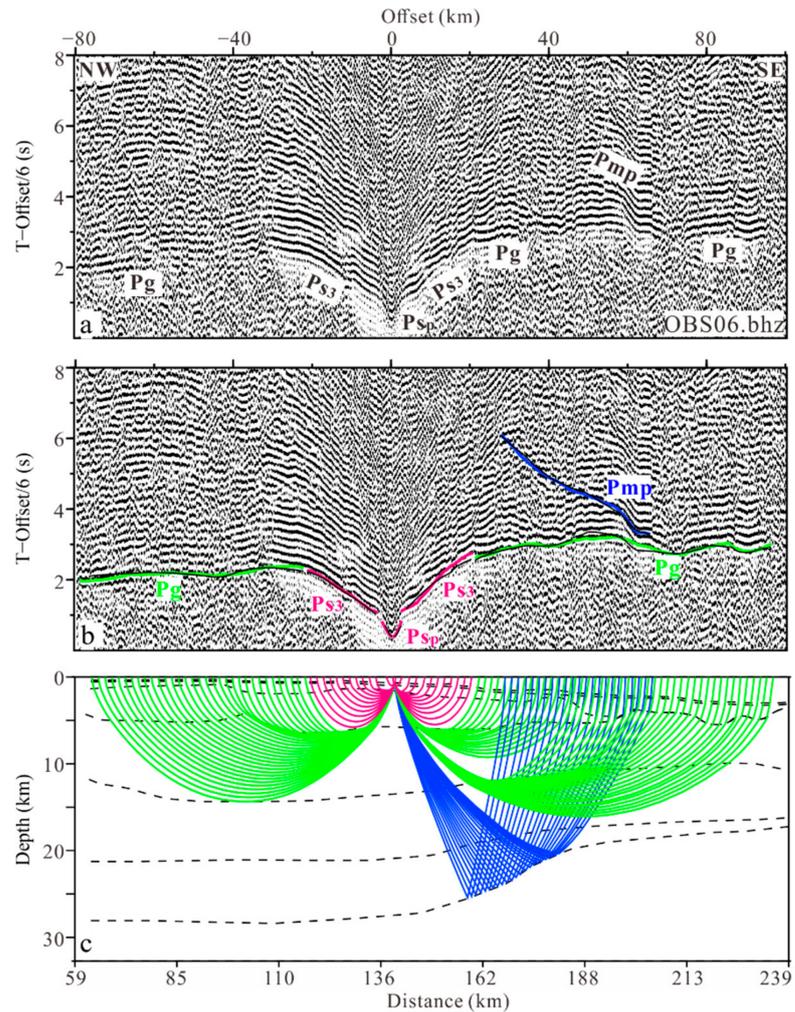


Figure 2. (a) Vertical seismic profile from OBS 06, with a reduction velocity of 6 km/s and a band-pass filter of 3–10 Hz (see Figure 1 for location). (b) The phase picking and fitting of OBS 06. The coloured lines represent picked travel times, and the thin, black lines represent calculated travel times from the final velocity model. (c) Ray tracing for OBS 06. Different line colours correspond to different phases.

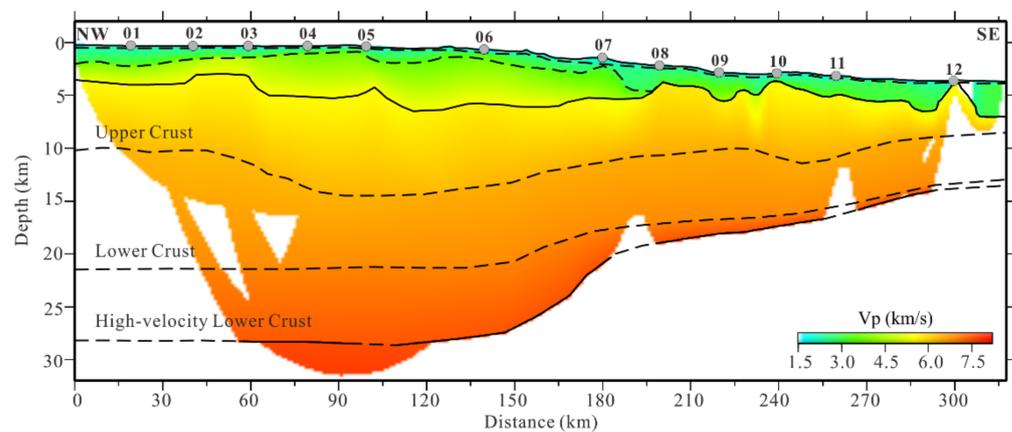


Figure 3. The forward P-wave velocity model along profile NS5. Gray circles represent OBS locations.

3. Types and Applications of Converted S-Waves Phases

In the OBS surveys, only the incident P-wave can reach the seafloor and propagate for the seismic waves that are generated by firing an air-gun in seawater. The S-wave phases recorded by the OBS are converted from the P-wave at different subsurface interfaces, such as the seafloor, sedimentary basement, and crust-mantle boundary [4,9,22]. The energy of converted S-waves depends on the angle of incidence [23]. The same parameters of filtering and reduction velocity were used to plot seismic profiles of the vertical and radial components. The converted S-wave phases were determined by comparing the kinematic (e.g., travel times, apparent velocity) and dynamic features (e.g., particle motion paths) of arrival phases in the vertical and radial components [4]. Compared to the vertical component, the converted S-wave phases in the radial component have lower apparent velocities and slower travel times (Figure 4a–d). The particle motion paths of the 710th trace in the OBS09 vertical component (Figure 4e) show that the energy of the vertical component (V) is strongest at 2.88–3.48 s, suggesting that particle motion was dominated by P-waves. The energy at 4.66–5.26 s in the radial component (R) is the largest (Figure 4f), denoting that this particle motion was dominated by S-waves. Two modes of S-wave phases [16] generally observed on OBSs are PPS and PSS phases (Figure 5). The apparent velocity of PPS arrivals is similar to corresponding P-waves, whereas the apparent velocity of PSS arrivals is much slower (Figure 6).

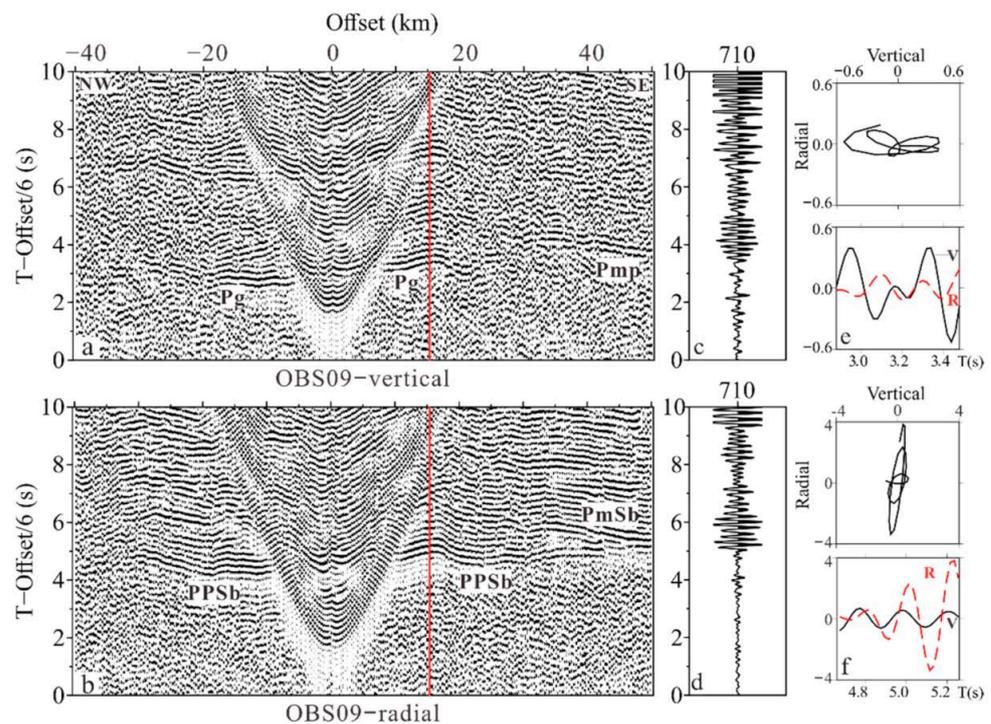


Figure 4. (a) Vertical and (b) radial component seismic profiles of OBS 09, reduced by 6 km/s. (c) and (d) show the waveform of the 710th trace in the vertical and radial component profiles, respectively. (e,f) correspond to the particle motion paths of the 710th trace in the vertical component (2.88–3.48 s) and radial component (4.66–5.26 s), respectively.

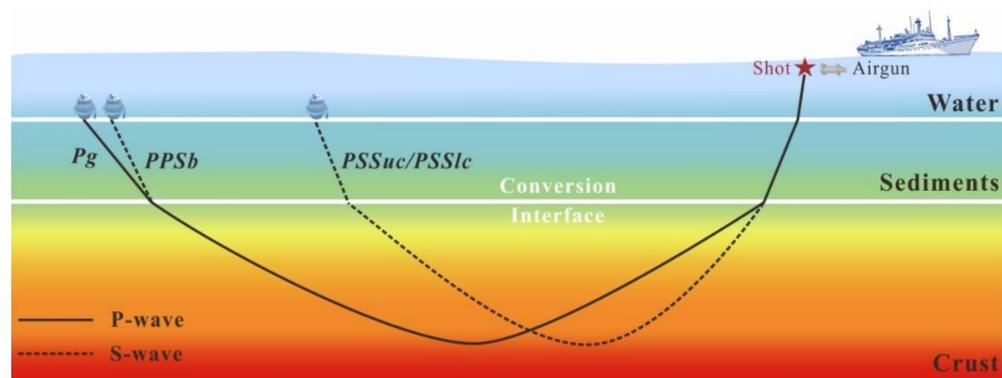


Figure 5. Cartoon illustrating the ray paths of converted S-wave phases. The PPS phases (PPSb) constrain the sediments S-wave velocity, while the PSS phases (PSSuc/PSSlc) constrain the crustal S-wave velocity.

3.1. PPS Phases and Its Role in Constraining Sediments V_p/V_s Ratios

The sediments' V_p/V_s ratios were constrained well by PPS phases, which were generated by P-waves converted to S-waves, at the interface, on the way up, and across the sediment as S-waves (Figures 5 and 6). The geometry of the S-wave model was kept consistent with the P-wave model, and the sediment interface depths were fixed and only the Poisson's ratio (σ) of different blocks was changed. The RayInvr software [17] was used to test the various Poisson's ratios within sedimentary layers until the best agreement was obtained between the calculated and observed PPS phase travel-times. The sediments S-wave velocity is calculated by the V_p/V_s ratios (the equation of V_p/V_s and Poisson's ratios is $\frac{v_p}{v_s} = \sqrt{1 + \frac{1}{1-2\sigma}}$) and the P-wave velocity.

The sedimentary layer in the P-wave model of the profile NS5 is divided into three layers (Figure 7), in which sedimentary layers 1 and 2 are Cenozoic sedimentary layers, and sedimentary layer 3 is Mesozoic strata. Sedimentary layer 1 is characterized by the thin thickness (0.16–0.5 km) and low V_p (1.7–2.2 km/s), hence the bottom of sedimentary layer 1 is a P- to S-wave conversion interface (e.g., PSSuc in Figure 6). Similarly, P- to S-wave conversion can also occur at the bottom of sedimentary layer 2, due to differences in petrophysical properties between sedimentary layer 2 and sedimentary layer 3 (e.g., PPSs2, PPSs3s2 in Figure 6). Refracted phases of Mesozoic sediments (P_s3) were converted to S-waves on the way up, at the bottom of sedimentary layer 2, and formed the PPSs3s2 phase. The PPSb phase was generated by the crustal refracted phases (P_g) converting to S-waves at the sedimentary basement on the way up, whereas the PPSs2 phase is from P_g phase converted at the bottom of sedimentary layer 2. The sediments' V_p/V_s ratios beneath the OBS were defined by the PPS phases. The S-wave velocity model of the sediment (Figure 7c) was obtained from dividing the previous P-wave velocity model (Figure 7a) by the V_p/V_s model (Figure 7b), which was calculated by the interpolation of the adjacent OBS.

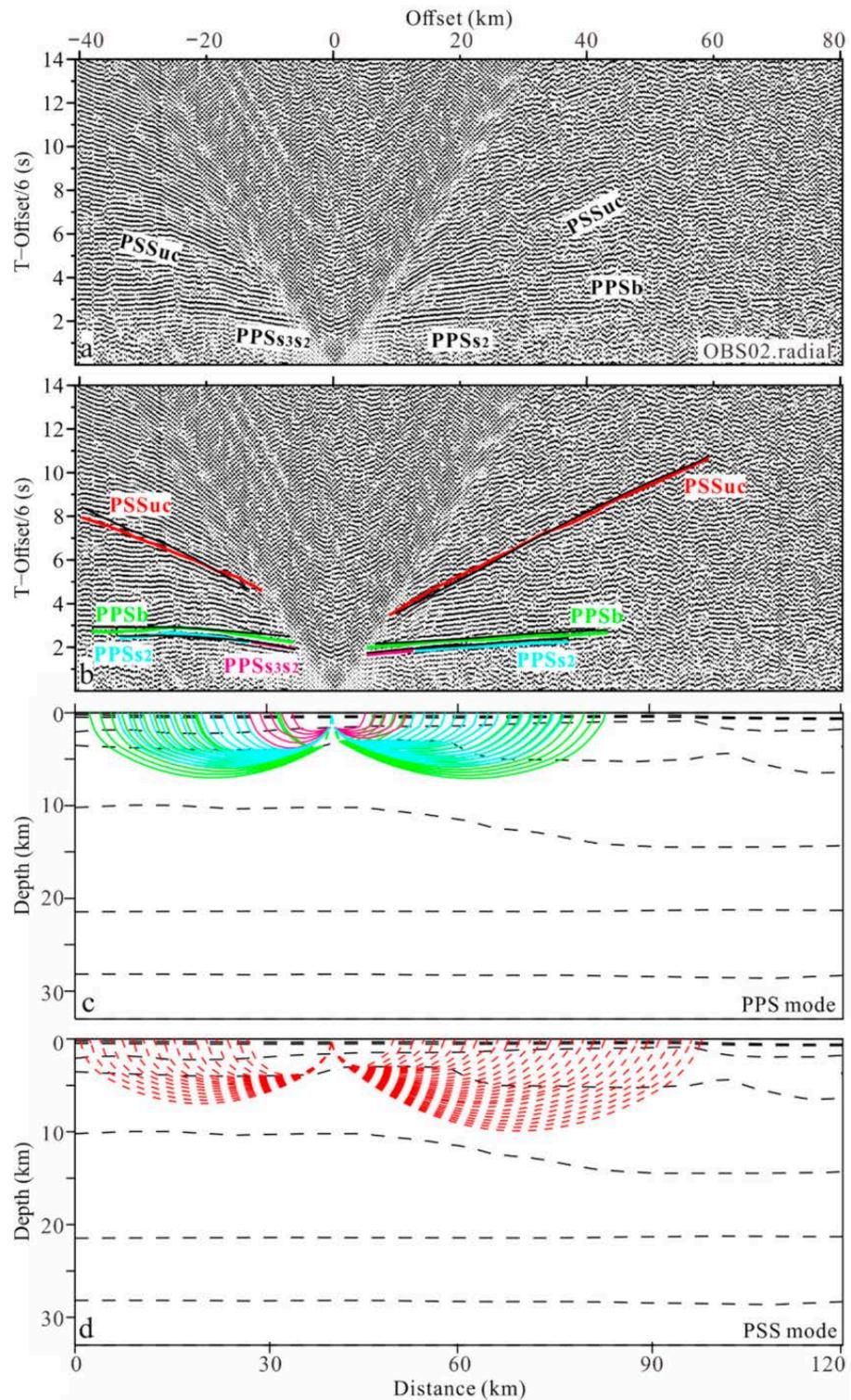


Figure 6. Seismic record section, S-wave phase identification, and ray tracing of OBS 02. (a) Seismic profile of radial component. (b) The identification and picking of S-wave phases. The coloured lines represent picked travel times and the thin black lines represent calculated travel times. (c,d) Show the ray tracing of PPS and PSS conversion modes, respectively. Solid lines represent P-wave paths, dotted lines represent S-wave paths, and the different coloured lines indicate different phases.

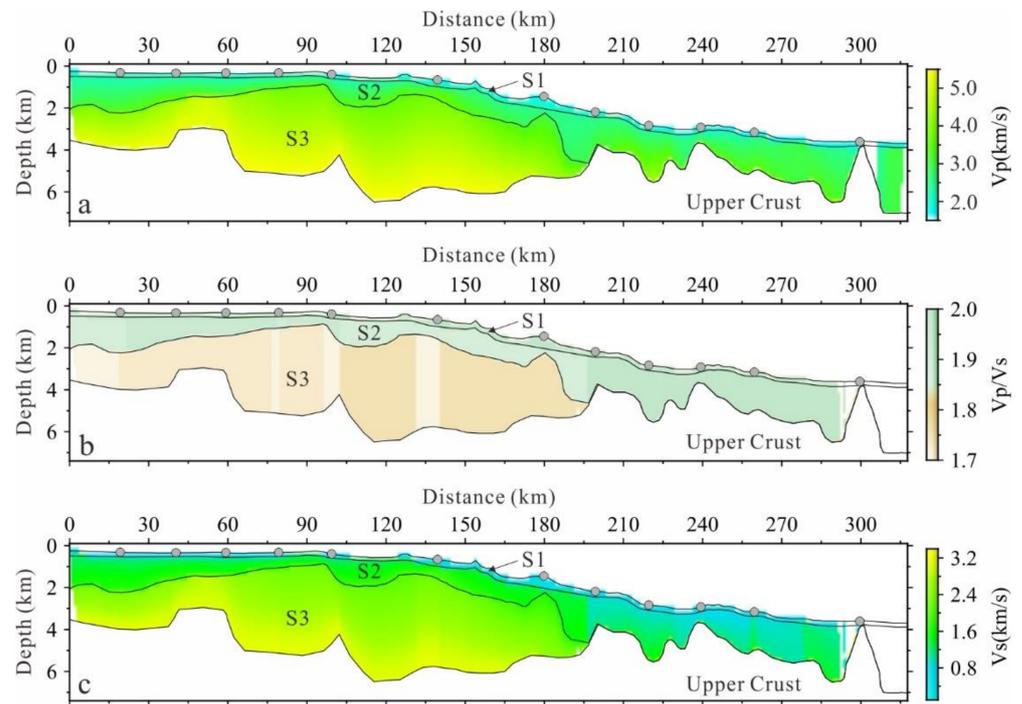


Figure 7. (a) P-wave velocity model, (b) V_p/V_s model, (c) S-wave velocity model of sedimentary layers. S1, S2 and S3 represent sedimentary layers 1, 2, and 3, respectively.

Our results show that sedimentary layer 1 has a high V_p/V_s of 3.05–7.14, representing unconsolidated sediments. The V_p/V_s of sedimentary layer 2 is also relatively high, with an average of 1.90 on the landward and 5.10 on the oceanward. The Mesozoic strata are characterised by high velocities and low V_p/V_s ratios of 1.71–1.76, due to the effects of long-term tectonism, sedimentary compactions, and later magmatism. The V_p/V_s of the sedimentary layers reveal two trends: one is that the V_p/V_s gradually decrease with the increased depth, which resulted from the confining pressure with depth; another is a seaward increase in V_p/V_s ratios, which may be related to composition and consolidation.

3.2. PSS Phases and Their Travel-Times Correction

The PSS phases result from P-waves are converted to S-waves at the sediment interface on the way down, so the entire path through the crust was at S-wave velocities (Figures 5 and 6). The PSSuc and PSSlc phases originated from the Pg phase refracted in the upper and lower crust, respectively, which were converted to S-waves at the bottom of sedimentary layer 1 on the way down (Figure 6). The PSSm phase was generated by Moho reflected phases (Pmp) converting to S-waves at the bottom of sedimentary layer 1 on the way down, while the Pmp phases converted at the basement formed the PSSmb phase.

The V_p/V_s ratios of the crust were constrained using the PSS phases recorded on the radial component. Since the converted interface of the PSS phases was within the sediments, its ray path included both P- and S-wave velocities segments. However, there is not a travel-time inversion method for the joint inversion of P- and S-wave velocities. The travel-time of the PSS arrivals can be corrected when the P-wave velocity and V_p/V_s ratios of the sedimentary layer are determined [24,25], because RayInvr allows the flexibility to model different arrival phases and conversions. Therefore, the conversion interface of the PSS phases was corrected from the sediments' interfaces to the seafloor. In that case, the velocity models have S-wave velocity below the seafloor, so the S-wave velocity inversion can be performed by the RayInvr software.

We verified the precision of the travel-time correction by ray-tracing the model using the corrected travel-times. The travel-time difference refers to the propagation time difference between the seafloor and the conversion interface at the velocities of P- and S-waves,

respectively. The fitting of the initial travel-time and corrected travel-time were compared by ray tracing the forward and composite model, which shows very small differences in the RMS. Corrected travel-times gave a fit with an RMS of 188 ms, compared to 185 ms for the initial travel-times (for details, see Table 1). In addition, the ray path of corrected travel-time is almost consistent with the initial (Figure 8), indicating that the correction of travel-time is reasonable.

Table 1. The total root-mean-square travel-time residuals before (RMS_b) and after (RMS_a) travel-times correction, picking uncertainties, conversion interface, for PSS mode arrival phases.

Stations	Phases	Uncertainties (ms)	RMS_b (ms)	RMS_a (ms)	Conversion Interface
OBS01	PSSuc	80	82	86	¹ S1
OBS02	PSSuc	180	182	183	¹ S1
OBS03	PSSuc	160	189	189	¹ S1
OBS05	PSSmb	200	181	179	Basement
OBS05	PSSm	180	150	141	¹ S1
OBS07	PSSuc	140	102	105	¹ S1
OBS07	PPSlc	140	111	111	¹ S1
OBS07	PSSm	180	221	221	¹ S1
OBS08	PSSuc	160	239	241	¹ S1
OBS08	PSSm	200	141	144	¹ S1
OBS09	PSSmb	200	185	167	Basement
OBS09	PSSuc	160	234	240	¹ S1
OBS10	PSSmb	160	197	198	Basement
OBS11	PSSuc	160	256	262	¹ S1
OBS12	PSSuc	200	193	199	¹ S1
OBS12	PPSlc	200	205	214	¹ S1

¹ S1: sedimentary layers 1.

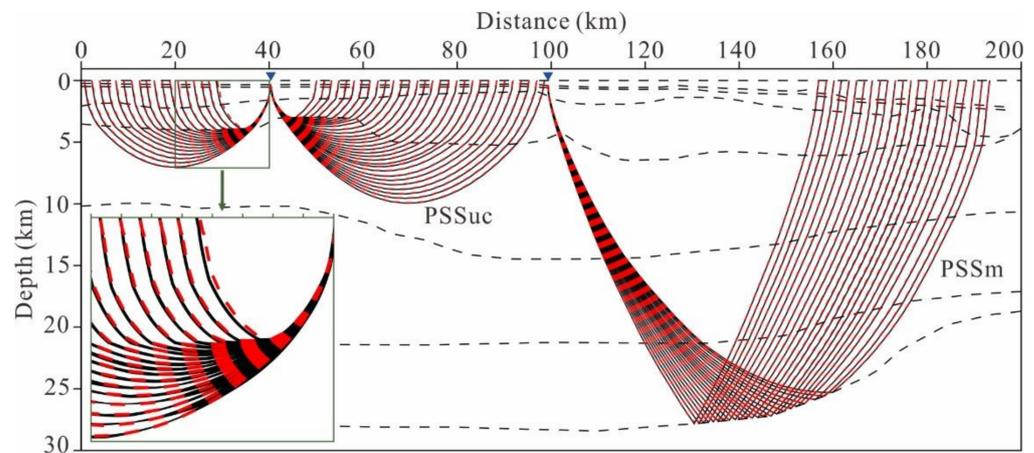


Figure 8. Comparison of the ray paths before and after the PSS phases' travel-time correction; The red dotted lines represent the original ray path, and the black solid lines represent the corrected ray path. The left-lower map shows a partial enlarged view of the 20–40 km in the model.

4. Crustal S-Wave Velocity Inversion

We only used the corrected PSS phases for the crustal S-wave velocity inversion, while the S-wave velocity of the sediments, constrained by the PPS phases, stays fixed during inversion. The RayInvr inversion was a process used to invert and update the velocity at individual nodes of the crust, the geometry of the model interfaces was kept consistent with the P-wave model, and the interface depths were kept constant. V_p/V_s ratios were very sensitive to the perturbation of the S-wave velocity model, which can supervise the instabilities during the inversion. The crustal V_p/V_s calculated from dividing the P-wave velocity model by the inverted S-wave velocity model was monitored during the inversion to ensure that the V_p/V_s ratios remained reasonable.

The damping factor (DF) will affect the convergence rate of the RMS. A small damping factor makes convergence faster, and the inversion result gradually approached

the least-squares solution, whereas the slower convergence rate is performed with a large damping factor. The numerical experiments show that a large damping factor will reduce the efficiency of the RayInvr algorithm, resulting in the number of iterations being increased and the resolution being reduced [26]. Selecting an appropriate damping factor can not only improve the inversion efficiency, but also make the inversion stable and reliable. After a series of tests, $DF = 3$ was finally selected as the damping factor for the following inversion. The velocity nodes of inversion are the same as those of P-wave model. The average inversion grid is set to 14, 16.5 and 19 km in the upper crust, lower crust and high-velocity lower crust, respectively (Figure S3).

4.1. Initial Models and the Iteration Termination Conditions

The initial S-wave velocity model is obtained from dividing the P-wave model by the set crustal V_p/V_s ratios. The five initial S-wave velocity models originated from starting V_p/V_s ratios of 1.71, 1.73, 1.75, 1.78, and 1.81 are shown in Figure 9. The V_p/V_s ratios varying from 1.71 to 1.81 represent the gradual change from felsic to mafic [27], which is reasonable for the geological features of the study area. Output V_p/V_s resulting from the division of P- and S-wave velocities are shown in Figure 9c,f,i,l,o. The best theoretical calculation model is achieved using a starting V_p/V_s ratio of 1.81, and the RMS is 187 ms. However, low V_p/V_s is abnormally (the blue area in Figure 9o) observed at the edge and 150 km of the model—such a model is unreasonable. The inversion is terminated after two iterations when the starting V_p/V_s is 1.73. Comprehensive analysis is done of the RMS and output V_p/V_s , and the best inversion model (Figure 9i) is produced from the starting V_p/V_s of 1.75 (Figure 9g). The final inversion model has a relatively good fit with the RMS of 193 ms, χ^2 of 1.718, and there is no local anomaly.

Generally, the optimum model is characterised by the RMS and χ^2 tended to 0 and 1, respectively. Therefore, the RMS and χ^2 should be taken into consideration if inversion is to be terminated. In addition to having a good travel-time fit, the final model should also trace as many rays as possible. If most of the rays are not traced, then this model should not be adopted. It is very important to establish the necessary conditions to terminate the inversion iteration, because this process requires manual interaction. After several iterations, the χ^2 has a value of <2 and the variation range is within 0.2, so the χ^2 is not taken into account here. We employed the number of tracing phases (NTP) and the RMS as the criteria for termination of inversion iterations (Figure 10). The NTP and RMS gradually decreased, as the adjustment of the model (sometimes the NTP will increase), and the variation range of the NTP is mainly within 30. After three to six iterations, the NTP sharply decreased with no obvious variation in the RMS (red dots in Figure 10); this means that the inversion can be terminated at this time.

4.2. Inversion Results and Analysis

The final inversion model shows (Figure 11a,c) that the upper crust and lower crust have a V_s of 3.0–3.7 km/s and 3.7–4.4 km/s, respectively. On the upper slope (at profile distance of 0–183 km), the V_p/V_s of the upper crust and lower crust are 1.71–1.75 and 1.72–1.75, respectively. Beneath the lower slope (at profile distance of 183–289 km), the upper crustal V_p/V_s is 1.75–1.79, and lower crustal V_p/V_s is 1.75–1.80. The high-velocity lower crust is also observed in our model that has a V_s of 4.0–4.4 km/s, and a V_p/V_s of 1.72–1.78. The final inversion model has a similar trend in the lateral velocities and V_p/V_s ratios with the forward model, due to the small difference with ΔV_s of 0.1 km/s and $\Delta V_p/V_s$ of 0.03 between the two models (Figure 12). The inversion model reveals more details in the crustal structures. For example, the high V_p/V_s ratios of the crust below two buried volcanos [19] may be related to magma conduits (Figure 11c). Therefore, the inversion results obtained by the RayInvr inversion algorithm are reliable and detailed for crustal structures.

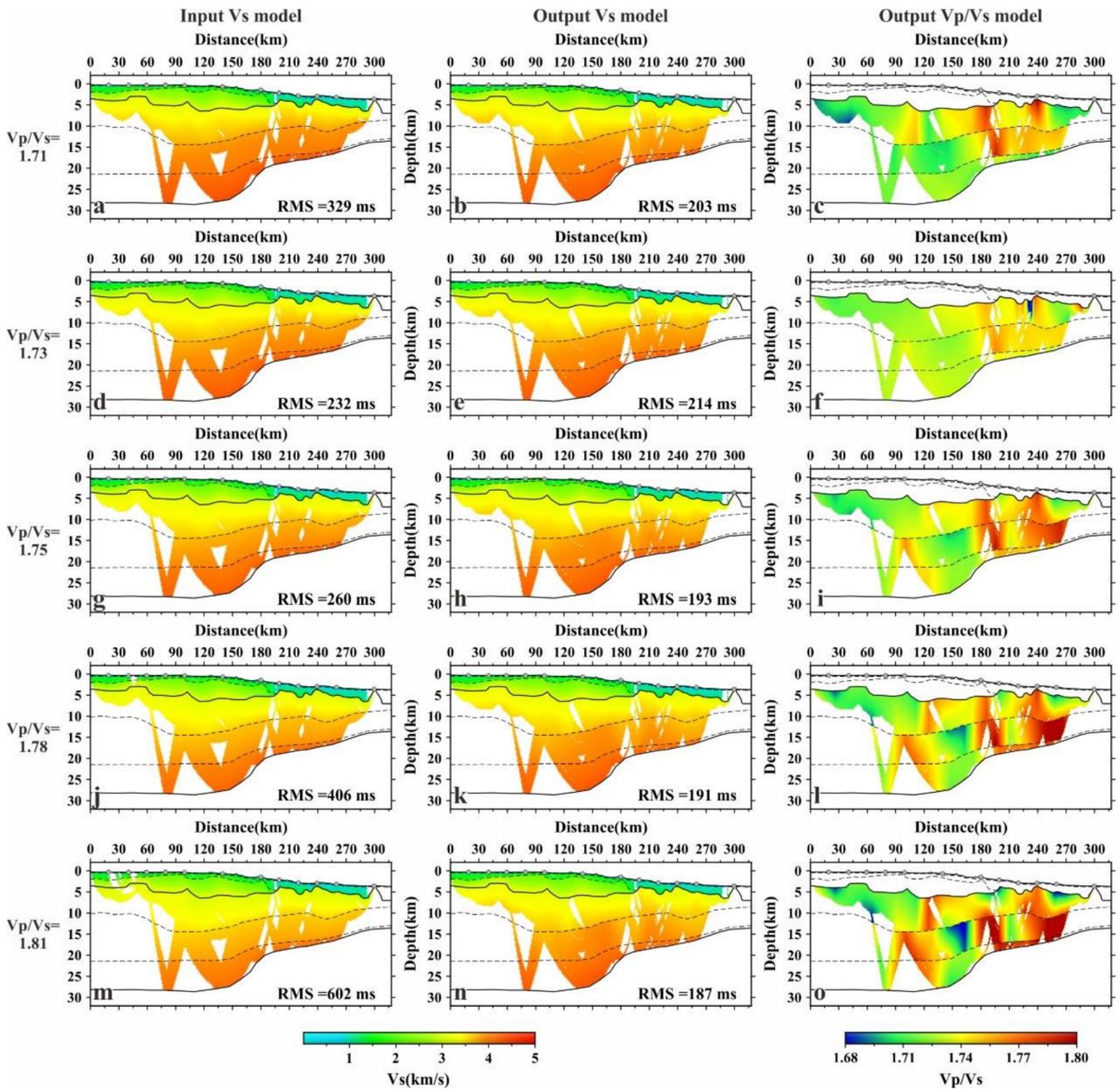


Figure 9. Inversion results for crustal S-wave velocity along profile NS5. The initial crustal S-wave velocity models of (a,d,g,j,m) in column 1 are originated from starting V_p/V_s ratios of 1.71, 1.73, 1.75, 1.78, and 1.81, respectively. Column 2 (b,e,h,k,n) shows the final results from inversion obtained from the different initial models in column 1 (a,d,g,j,m). Column 3 (c,f,i,l,o) shows apparent crustal V_p/V_s models from the division of the P-wave velocity model by the final S-wave velocity models in column 2 (b,e,h,k,n).

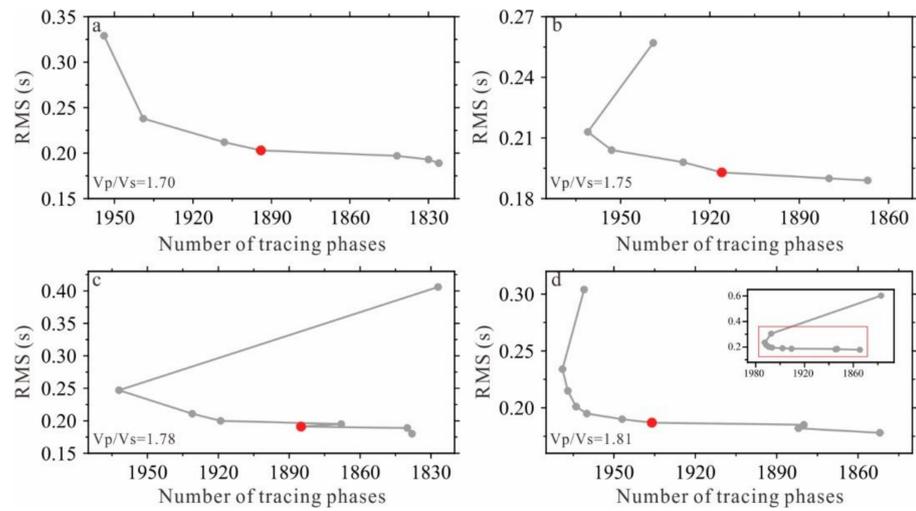


Figure 10. Variation in the number of tracing phases and travel-time residuals (RMS) during the inversion process under the initial models with the V_p/V_s ratios of 1.71 (a), 1.75 (b), 1.78 (c), 1.81 (d). Red dots indicate inversion is terminated at this time.

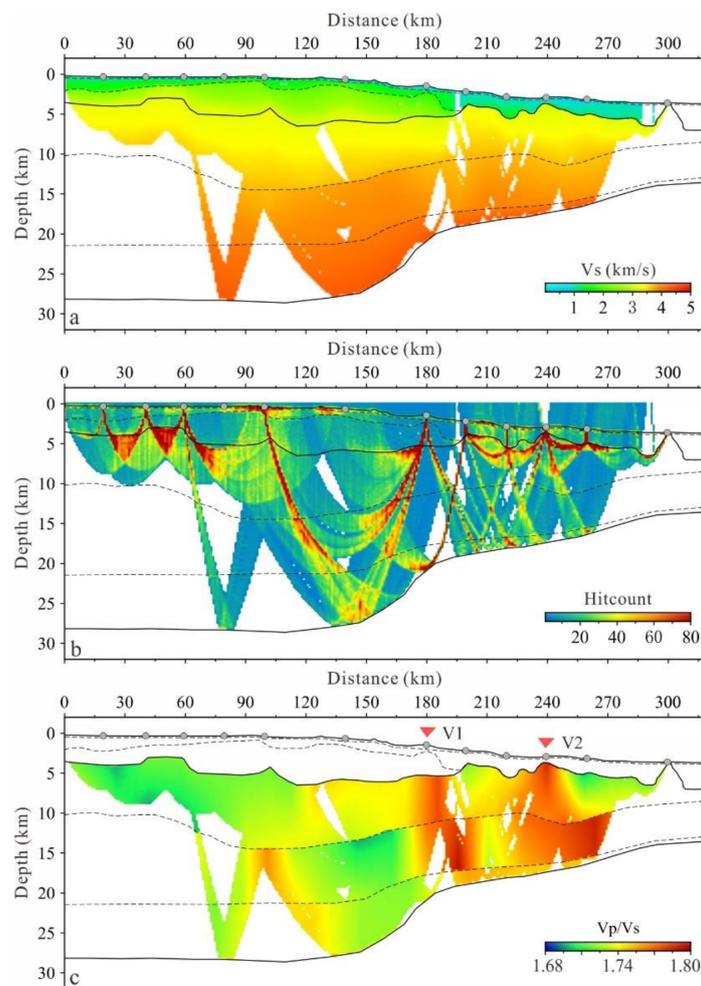


Figure 11. (a) The final inversion result for crustal S-wave velocity model. (b) Ray coverage density distribution, and the dimensions of counting grid are 1×0.25 km. (c) Apparent crustal V_p/V_s model that calculated by dividing the P-wave model by the inverted S-wave model (a). V1, V2 represent the buried volcanos [19].

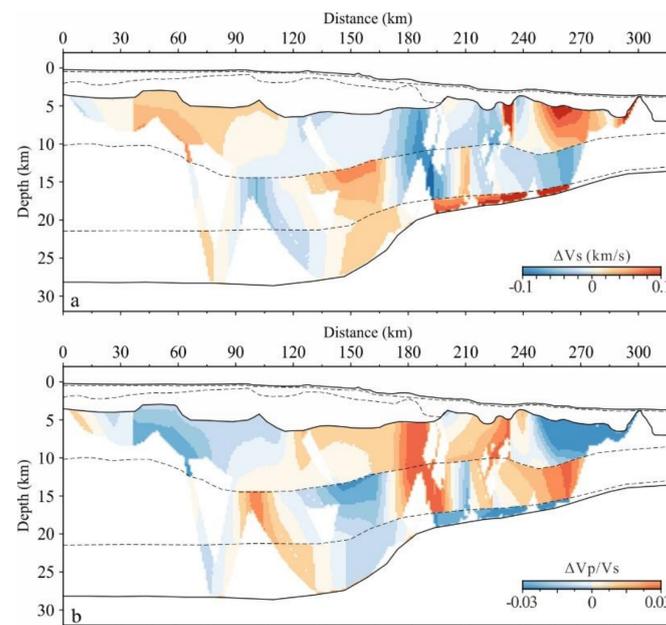


Figure 12. The difference between the forward and inversion results for crustal S-wave velocity (a), V_p/V_s ratios (b). ΔV_s and $\Delta V_p/V_s$ are the difference between inversion and forward results.

Figure 11b shows that the model has a good ray coverage with the ray coverage of most areas exceeding 20–40 hits. To better determine the resolution of the inversion model, we imposed a checkerboard pattern on the model [28]. This grid pattern had a width of 15–25 km in the upper crust, 20–40 km in the lower crust, and the vertical dimension was constrained by the model interface. Furthermore, we added S-wave velocity perturbations of ± 0.2 km/s that were equal in amplitude but opposite in terms of the sign at adjacent nodes (Figure 13a). The output model is shown in Figure 13b. The checkerboard test confirmed the good resolutions of the lower crust, especially in the upper crust.

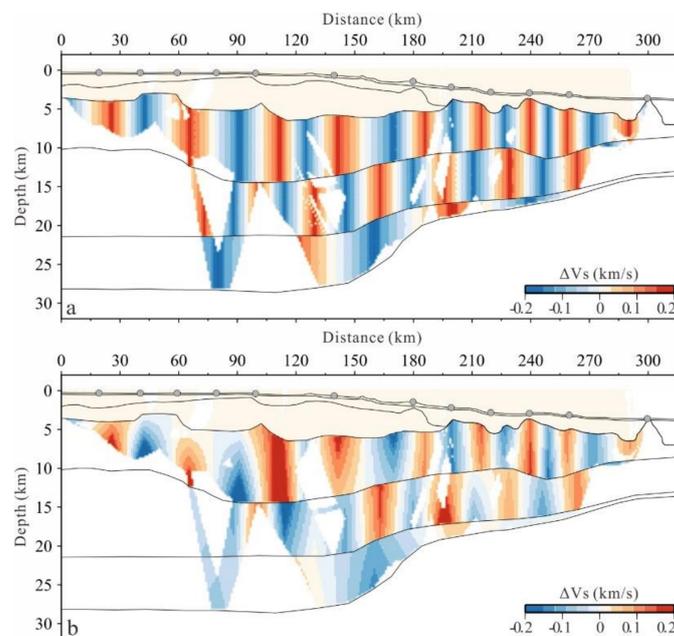


Figure 13. (a) Synthetic S-wave velocity perturbation using a checkerboard test. (b) Calculated S-wave velocity perturbation for the checkerboard test. The horizontal size of the anomalous block was 15–25 km in the upper crust, 20–40 km in the lower crust, the vertical size was bounded by interface.

5. Conclusions

We skillfully performed the travel-time inversion of converted S-waves based on RayInvr algorithms by transforming the conversion interface of the sedimentary layers to the seafloor. The iterative inversion is terminated by reaching the conditions of the travel-time residuals (RMS) and the number of tracing phases (NTP). The P-wave velocity model and the inverted S-wave velocity mode are presented to calculate the crustal V_p/V_s ratios. Our results show the crust has a V_s of 3.0–4.4 km/s, and the crustal V_p/V_s ratios are 1.71–1.80, which are reliable due to the small difference, with a ΔV_s of 0.1 km/s and $\Delta V_p/V_s$ of 0.03, between the inversion and forward model. In addition, the high-resolution inversion model has revealed many details of the crustal structures, including magma conduits.

Supplementary Materials: The following are available online at <https://www.mdpi.com/article/10.3390/app11083571/s1>, Figure S1: Results of the forward P-wave velocity, Figure S2: The F-test results for the uncertainty of the Moho depth and velocity in the layer of high-velocity lower crust, Figure S3: The velocity grid nodes for the inversion, Table S1: The polarization angle for each OBS.

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