Abstract: This contribution reports on the observation of enigmatic fibrous platinum-group minerals (PGM) found within a chromitite body included in limonite (“floating chromitite”) from Ni-laterites in the Dominican Republic. Fibrous PGM have a Ru-Os-Ir-Fe dominated composition and are characterized by fibrous textures explained by grain-forming fibers which are significantly longer (1–5 µm) than they are wide (~100 nm). Back-scattered electron (BSE) images suggest that these nanofibers are platinum-group elements (PGE)-bearing and form <5 µm thick layers of bundles which are oriented orthogonal to grains’ surfaces. Trace amounts of Si are most likely associated with PGE-bearing nanofibers. One characteristic fibrous PGM was studied in detail: XRD analyses point to ruthenian hexaferrum. However, the unpolished fibrous PGM shows numerous complex textures on its surface which are suggestive for neoformation processes: (i) features suggesting growth of PGE-bearing nanofibers; (ii) occurrence of PGM nanoparticles within film material (biofilm?) associated with PGE-bearing nanofibers; (iii) a Si-rich and crater-like texture hosting PGM nanoparticles and an Ir-rich accumulation of irregular shape; (iv) complex PGM nanoparticles with ragged morphologies, resembling sponge spicules and (v) oval forms (<1 µm in diameter) with included PGM nanoparticles, similar to those observed in experiments with PGE-reducing bacteria. Fibrous PGM found in the limonite may have formed due to supergene (bio-)weathering of fibrous Mg-silicates which were incorporated into desulphurized laurite during stages of serpentinization.

Keywords: PGE; PGM; fibrous; nanofibers; Ni-laterite; supergene; weathering; neoformation; Falcondo; Dominican Republic

1. Introduction

The platinum-group elements (PGE) are a group of precious metals (Os, Ir, Ru, Rh, Pt, and Pd) which are referred to strategic metals because of their extreme scarcity within the Earth’s crust combined with the increasing need for PGE in advanced technologies (e.g., autocatalyst industries; electronics) and jewellery [1]. Currently, the vast majority (~80%) of the worldwide PGE demand is supplied by magmatic ore deposits located in only two countries, i.e., South Africa and Russia [2]. This supply situation has motivated the search for alternative sources of PGE, and large oxidized ore deposits associated with near surface modification of ultramafic rocks, such as ophiolite-related Ni-laterites, have been considered as worthy exploration targets for future unconventional PGE ore deposits.
deposits [3]. In general, total PGE contents found in Ni-laterites around the world are in the range of <100 ppb to up to a few hundred ppb, however lateritic crust with 2 ppm PGE was reported from Western Australia [4] as well as >4 ppm PGE from Burundi [5].

Ni-laterites from the northern Caribbean are systematically enriched in the limonite and reach values of up to 0.6 ppm total PGE [6]. These increasing PGE contents toward the surface are mainly explained by residual accumulations of PGE-bearing minerals. However, the results of detailed mineralogical surveys also provide convincing morphological and chemical evidence for in situ neoformation of platinum-group minerals (PGM) within Ni-laterites by bioreduction and/or electrochemical accretion [7,8]. Hence, accumulations of what are most likely biogenic mediated in situ growths of PGM nanoparticles are believed to contribute to PGE anomalies in Ni-laterites. The recent verification of the biological role of PGM transformations in supergene environments by [9] has marked the beginning of a new period of PGE-related investigation and exploration.

The presented contribution reports on the observation of enigmatic fibrous PGM found close to the surface in Ni-laterites from the Dominican Republic and aims (i) to give a general overview of the prevailing features of fibrous PGM; (ii) to provide detailed textural as well as mineralogical data of one characteristic fibrous PGM and (iii) to discuss a possible origin of fibrous PGM.

2. Sample Material and Methods

Samples were collected in the ophiolite-related Loma Larga Ni-laterite ore deposit which forms part of the Falcondo mining area located in the central part of the Dominican Republic (Figure 1). The Falcondo mining area is currently known as the largest hydrous Mg-silicate type Ni-laterite deposit of the Caribbean, with measured and indicated Ni resources of 67.8 Mt at a grade of 1.5% Ni plus 4.9 Mt at 1.4% Ni inferred [10].

The Ni-laterite at Falcondo developed over the so-called Loma Caribe peridotite which occurs as a serpentinized belt of ultramafic rocks approximately 4–5 km wide and 95 km long in the northwest of Santo Domingo (Figure 1). Detailed information of the local geology and formation history of the Ni-laterites can be found elsewhere (e.g., [11] and references therein). In general, the weathering profile of the Falcondo mining area is divided into: (i) serpentinized protolith (harzburgite > dunite > lherzolite) at the bottom followed by (ii) saprolite with hydrous Mg-silicate dominated mineralogy (beneath the Mg-discontinuity) and (iii) limonite with Fe-oxide(s) dominated mineralogy at the top (above the Mg-discontinuity) [11–14].

Small, rare, and hence uneconomic chromitite occurrences are scattered within given horizons of the lateritic profile [15,16]. Within the Ni-laterite profile of the Loma Larga ore deposit a highly weathered chromitite body of approximately 0.5 m × 0.5 m is included in limonite (“floating chromitite”) (Figure 1). This chromitite is characterized by extremely high PGE contents up to 17.5 ppm total PGE [7].

Approximately 100 g of homogenized “floating chromitite” sample was carefully disaggregated, milled, and hand-sieved before being processed by means of hydroseparation (HS) techniques at the HS laboratory in Barcelona, Spain ([8] and reference therein). Heavy mineral concentrates were mounted as loose (unpolished) grains on a metallic cylinder (diameter 2.5 cm) and/or as polished monolayer resin blocks and were subsequently investigated by stereomicroscopy, optical reflected light microscopy, and by scanning electron microscopy (SEM) using a Quanta 200 FEI XTE 325/D8395 (FEI, Hillsboro, OR, USA) as well as a field emission scanning electron microscope (FE-SEM) Jeol JSM-7100 (JEOL, Tokyo, Japan) at the Serveis Científics i Tecnòlogics, University of Barcelona, Spain.

At the same institution, quantitative electron microprobe analyses (EMPA) were obtained with a JEOL JXA-8230 electron microprobe (JEOL, Tokyo, Japan) operating in wavelength-dispersive spectroscopy (WDS) mode with an accelerating voltage of 20 kV, a beam current of 10 nA, and a beam diameter of 1 µm. Pure metals were used as standards for Os, Ir, Ru, Rh, Pt, Pd, Co, and Sb as well as chromite (Al, Fe), periclase (Mg), NiO (Ni), S (S), GaAs (As), and wollastonite (Si). The following interferences RuLβ → RhLα, RhLβ → PdLα, RuLβ → PdLα, and RhLα → PtLα were online-corrected.
The same instrument was used for element distribution maps (WDS mode) with an accelerating voltage of 20 kV and a beam current of 128.8 nA. Maps were collected by beam scanning with dwell times of 60 ms/pixel. For each element, the background map was subtracted from the corresponding peak map.

Figure 1. (A) The location of the Loma Caribe peridotite and an orthophotograph of the Falcondo mining area highlighting the Loma Larga ore deposit. The inset shows an idealized Ni-laterite soil profile from the Falcondo mining area as well as a field observation of the PGE-rich “floating chromitite” within limonite (width of image: ~1 m); (B) Simplified geological map of the central section of the Loma Caribe peridotite (modified from [17,18]).

One fibrous PGM was investigated by single-crystal X-ray diffraction (XRD, Bruker, Billerica, MA, USA) analysis at the same institution. The X-ray intensity data were measured on a D8 Venture system equipped with a multilayer monochromator and a Mo microfocus ($\lambda = 0.71073$ Å). The sample was analyzed at a distance of 34 mm. The angles of the goniometer were $\theta = 0^\circ$, $\omega = 360^\circ$, and $\chi = 90^\circ$, and the sample was spun $360^\circ$. The measuring time was 60 s per step. Subsequent results were processed with Bruker software [19] which was used to subtract the background of the patterns, to detect the peaks, and to assign minerals and their corresponding $d_{hkl}$ to each peak.

3. Mineralogy of Fibrous PGM

Approximately 300 grains of PGM in the size range of <10 µm to ~125 µm in diameter and with either Ru-, Ir-, or Pt-dominated mineralogy were detected in heavy mineral concentrates. Ru-dominated and Fe-rich alloys (e.g., ruthenian hexaferrum) are the most abundant PGM (>90%), followed by lower quantities of Ir- and Pt-bearing Fe-Ni alloys ([7,20] respectively). Careful examination of all PGM via FE-SEM led to the observation of a rarely occurring (~1%), Ru-dominated subgroup which is characterized by fibrous textures (Figure 2). These textures are explained by grain-forming fibers which are significantly longer (1–5 µm) than they are wide (~100 nm). Due to their nanometric diameters, these fibers were defined as nanofibers. In general they occur as <5 µm thick layers of nanofiber-bundles which are oriented orthogonal to the grains’ surfaces (Figure 2B,E). Fibrous PGM show diverse morphologies that range from subhedral to anhedral grains (e.g., Figure 2C,I, respectively) and elongated fibrous PGM occur in addition (Figure 2G). Interestingly, some grains contain rounded voids from where nanofiber-bundles propagate radially (Figure 2D,E,I). Energy-dispersive spectroscopy (EDS) analyses (spot diameter of 1 µm) suggest that fibrous PGM contain mainly Ru and Fe with minor contents of Os and Ir. However, small but clear peaks of Si are also observed.
3.1. Detailed Investigation of One Fibrous PGM

3.1.1. Field Emission Scanning Electron Microscopy (FE-SEM)

One fibrous PGM was separated and investigated in detail. The unpolished grain has an isometric morphology with approximately 75 µm long axes and is characterized by a subhedral shape with damaged corners and a broken central part (Figure 3A). A detailed survey by means of FE-SEM revealed numerous features on the surface of the grain which, to the best of the authors’ knowledge, have not been described for PGM before:

- the structure of the grain suggests an interdependency of PGM nanoparticles and nanofiber-bundles which are oriented orthogonal to the grain’s surface (Figure 3A,B,G,H);
- nanofibers occur within layers of nanofiber-bundles as well as within open spaces suggesting either in situ growth (“healing of cracks”) or misalignments (Figure 3C,D);
- comparison of BSE images and secondary electron (SE) images suggests that elements of high atomic number (PGE) are present in the central part of the nanofibers (bright in BSE images),
whereas the outer part of the nanofibers appears to consist of elements with lower atomic numbers (e.g., Si) (Figure 3E,F);

- at the endings of nanofiber-bundles, a thin layer of a homogenous film occurs that host PGM nanoparticles (Figure 3G–J);
- within the central part of the grain a Si-rich, crater-like texture is present which occurs dark in BSE images (Figure 3A). This “crater” measures approximately 5 µm × 7 µm in diameter and hosts a great number of homogenously distributed PGM nanoparticles. At its periphery an Ir-rich accumulation (~1 µm in diameter) with a rather irregular shape is observed (Figure 3K);
- around the crater-like texture, PGM nanoparticles occur, which are characterized by ragged morphologies resembling sponge spicules. Interestingly, the Si-rich crater-like texture hosts smaller PGM nanoparticles than its Si-depleted immediate vicinity (Figure 3K);
- on plain walls at the upper part of the grain (Figure 3A), which represent the backside of nanofiber-bundles and therefore most likely film material, oval forms (~500 nm at the long axes) with included PGM nanoparticles occur, resembling observations by [21] in experiments with PGE reducing bacteria (Figure 3L).
Figure 3. BSE and secondary electron images (SE) of a fibrous PGM showing numerous features: (A) overview; note the isometric, subhedral shape of the grain; (B) close-up of the platinum-group elements (PGE)-bearing nanofiber-bundles; (C) features suggesting either “healing of cracks” or misalignments of nanofibers; (D) layers of PGE-bearing nanofiber-bundles which are oriented orthogonal to the grain’s surface; (E) SE image close-up of nanofibers; (F) BSE image close-up of the same area as in (E); note the bright occurring center of the nanofibers in the BSE image, suggesting the presence of PGE; (G,H) close-up of the contact between PGE-bearing nanofiber-bundles and associated film material (biofilm?) hosting PGM nanoparticles in SE (G) and BSE (H); (I,J) close-up of the film material showing layering and numerous PGM nanoparticles in SE (I) and in BSE (J); (K) detail of the zone around the crater-like texture in the central part of the grain with associated Ir-rich accumulation and complex morphologies of PGM nanoparticles; note the different size and shape of PGM nanoparticles within the dark (Si-rich) crater-like texture compared to its immediate vicinity; (L) approximately 500 nm long oval forms with included PGM nanoparticles resembling observations related to PGE depositing bacteria [21].

3.1.2. X-ray Diffraction (XRD)

The same grain as shown in Figure 3A, revealing metallic luster under the stereomicroscope (Figure 4A), was studied using a single-crystal X-ray diffractometer. The obtained diffraction pattern displayed characteristic concentric, well defined, and homogeneous Debye rings, indicating that the
grain is constituted by a randomly oriented polycrystalline assemblage (Figure 4B). Integration of the intensity data at constant steps of 2θ gives a conventional one dimensional powder diffraction pattern (Figure 4C).

A comparison with literature data for chemically related mineral phases (Os-group minerals) revealed that the fibrous PGM has rather identical X-ray line intensities and observed d-spacings as ruthenian hexaferrum [20] (Table 1). On this basis, the X-ray powder pattern for the fibrous PGM was indexed, leading to the refined unit-cell dimensions: \(a = 2.6670\) (13) \(Å\), \(c = 4.2361\) (18) \(Å\), \(V = 26.095\) (27) \(Å^3\), \(Z = 2\) (Table 2).

Table 1. Powder X-ray data for the fibrous platinum-group mineral (PGM) shown in Figure 3A and comparison with chemically related phases.

<table>
<thead>
<tr>
<th>Fibrous PGM</th>
<th>Ruthenian Hexaferrum (^1)</th>
<th>Hexaferrum (PDF 54-0704)</th>
<th>Ruthenium (PDF 00-06-0663)</th>
<th>Osmium (PDF 00-041-0601)</th>
</tr>
</thead>
<tbody>
<tr>
<td>(hkl)</td>
<td>(d)</td>
<td>(hkl)</td>
<td>(d)</td>
<td>(hkl)</td>
</tr>
<tr>
<td>20</td>
<td>2.3097</td>
<td>100</td>
<td>20</td>
<td>2.3068</td>
</tr>
<tr>
<td>25</td>
<td>2.1161</td>
<td>002</td>
<td>20</td>
<td>2.1250</td>
</tr>
<tr>
<td>100</td>
<td>2.0279</td>
<td>101</td>
<td>100</td>
<td>2.0274</td>
</tr>
<tr>
<td>30</td>
<td>1.5611</td>
<td>102</td>
<td>25</td>
<td>1.5629</td>
</tr>
<tr>
<td>35</td>
<td>1.3335</td>
<td>110</td>
<td>40</td>
<td>1.3318</td>
</tr>
<tr>
<td>40</td>
<td>1.2047</td>
<td>103</td>
<td>50</td>
<td>1.2072</td>
</tr>
<tr>
<td>50</td>
<td>1.1285</td>
<td>112</td>
<td>60</td>
<td>1.1285</td>
</tr>
<tr>
<td>35</td>
<td>1.1142</td>
<td>201</td>
<td>40</td>
<td>1.1131</td>
</tr>
</tbody>
</table>

1 [20].

Table 2. Comparative data for chemically related phases including the fibrous PGM.

<table>
<thead>
<tr>
<th>Chem. Formula</th>
<th>Ruthenian hexaferrum (^1)</th>
<th>Hexaferrum (^2)</th>
<th>Ruthenium (^3)</th>
<th>Osmium (^4)</th>
<th>Garutite (^5)</th>
</tr>
</thead>
<tbody>
<tr>
<td>(Ru,Os,Ir)(<em>{10})Fe(</em>{5})</td>
<td>P6(_3)/mmc</td>
<td>P6(_3)/mmc</td>
<td>P6(_3)/mmc</td>
<td>P6(_3)/mmc</td>
<td>P6(_3)/mmc</td>
</tr>
<tr>
<td>(a)</td>
<td>2.6670 (13)</td>
<td>2.664 (1)</td>
<td>2.64 (1)</td>
<td>2.64 (1)</td>
<td>2.64 (1)</td>
</tr>
<tr>
<td>(c)</td>
<td>4.2361 (18)</td>
<td>4.2361 (18)</td>
<td>4.20 (2)</td>
<td>4.20 (2)</td>
<td>4.20 (2)</td>
</tr>
<tr>
<td>(V)</td>
<td>26.095 (27)</td>
<td>26.125 (3)</td>
<td>25.35</td>
<td>27.84</td>
<td>27.84</td>
</tr>
<tr>
<td>(Z)</td>
<td>2</td>
<td>2</td>
<td>2</td>
<td>2</td>
<td>2</td>
</tr>
</tbody>
</table>

1 [20]; 2 [22]; 3 [23]; 4 [24]; 5 [25].
A close-up image of the Si-bearing belt displays the presence of PGE-bearing nanofibers (Figure 5B). Quantitative analyses as well as for element mappings. Polishing of the fibrous PGM led to a general of slightly higher contents, whereas Ru does not show visible compositional variations (Figure 5C). However, within the Si-rich area, lower Fe and Ni contents are observed, Os reveals an elongated zone in Figure 5A confirm the above-mentioned dominance of Fe-rich, Ru-Os-Ir compositions (Figure 5C). To the possible occurrence of O or OH within the nanofibers. Element mappings of the area indicated in Figure 5A confirm the above-mentioned dominance of Fe-rich, Ru-Os-Ir compositions (Figure 5C). Low totals (<85.55 wt %) are mainly attributed to the fine-grained nature of the grain and maybe also to the possible occurrence of O or OH within the nanofibers. Element mappings of the area indicated in Figure 5A confirm the above-mentioned dominance of Fe-rich, Ru-Os-Ir compositions (Figure 5C). However, within the Si-rich area, lower Fe and Ni contents are observed, Os reveals an elongated zone of slightly higher contents, whereas Ru does not show visible compositional variations (Figure 5C).

### 3.1.3. Electron Microprobe Analysis (EMPA)

In continuation, the same grain as shown in Figure 3A was included in epoxy and polished for quantitative analyses as well as for element mappings. Polishing of the fibrous PGM led to a general loss of textures, providing evidence that observed complex textures occur exclusively associated with the surface of fibrous PGM. BSE images of the polished fibrous PGM reveal rather homogenous and bright zones, which are crosscut by a slightly darker, Si-bearing belt of varying thickness (Figure 5A). A close-up image of the Si-bearing belt displays the presence of PGE-bearing nanofibers (Figure 5B).

EMP analyses of four points (indicated in Figure 5A) are given in Table 3 and confirm the dominance of Ru, Os, Fe, and Ir relative to other elements including trace amounts (<1 wt %) of Mn, Si, Al, and Ni. Low totals (<85.55 wt %) are mainly attributed to the fine-grained nature of the grain and maybe also to the possible occurrence of O or OH within the nanofibers. Element mappings of the area indicated in Figure 5A confirm the above-mentioned dominance of Fe-rich, Ru-Os-Ir compositions (Figure 5C). However, within the Si-rich area, lower Fe and Ni contents are observed, Os reveals an elongated zone of slightly higher contents, whereas Ru does not show visible compositional variations (Figure 5C).

**Table 3.** Composition of the fibrous PGM analyzed by electron microprobe (EMP).

<table>
<thead>
<tr>
<th>S</th>
<th>As</th>
<th>Os</th>
<th>Ir</th>
<th>Ru</th>
<th>Rh</th>
<th>Pt</th>
<th>Pd</th>
<th>Fe</th>
<th>Ni</th>
<th>Cu</th>
<th>Co</th>
<th>Si</th>
<th>Al</th>
<th>Mg</th>
<th>Sb</th>
<th>Mn</th>
<th>Total</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>0.04</td>
<td>0.15</td>
<td>17.95</td>
<td>5.32</td>
<td>42.96</td>
<td>bdl</td>
<td>bdl</td>
<td>bdl</td>
<td>16.64</td>
<td>0.33</td>
<td>0.09</td>
<td>0.13</td>
<td>0.54</td>
<td>0.62</td>
<td>0.05</td>
<td>bdl</td>
<td>0.59</td>
</tr>
<tr>
<td>2</td>
<td>0.12</td>
<td>0.15</td>
<td>17.78</td>
<td>5.04</td>
<td>40.94</td>
<td>0.12</td>
<td>bdl</td>
<td>bdl</td>
<td>11.36</td>
<td>0.17</td>
<td>0.09</td>
<td>0.07</td>
<td>0.82</td>
<td>0.44</td>
<td>0.03</td>
<td>0.15</td>
<td>0.86</td>
</tr>
<tr>
<td>3</td>
<td>0.06</td>
<td>0.10</td>
<td>16.34</td>
<td>7.78</td>
<td>43.03</td>
<td>bdl</td>
<td>bdl</td>
<td>bdl</td>
<td>14.59</td>
<td>0.29</td>
<td>0.06</td>
<td>0.14</td>
<td>0.47</td>
<td>0.59</td>
<td>0.06</td>
<td>bdl</td>
<td>0.76</td>
</tr>
<tr>
<td>4</td>
<td>0.03</td>
<td>0.09</td>
<td>17.31</td>
<td>7.81</td>
<td>41.56</td>
<td>bdl</td>
<td>bdl</td>
<td>bdl</td>
<td>16.50</td>
<td>0.36</td>
<td>0.10</td>
<td>0.10</td>
<td>0.47</td>
<td>0.60</td>
<td>0.04</td>
<td>bdl</td>
<td>0.58</td>
</tr>
</tbody>
</table>

<table>
<thead>
<tr>
<th>S</th>
<th>As</th>
<th>Os</th>
<th>Ir</th>
<th>Ru</th>
<th>Rh</th>
<th>Pt</th>
<th>Pd</th>
<th>Fe</th>
<th>Ni</th>
<th>Cu</th>
<th>Co</th>
<th>Si</th>
<th>Al</th>
<th>Mg</th>
<th>Sb</th>
<th>Mn</th>
</tr>
</thead>
<tbody>
<tr>
<td>1</td>
<td>0.14</td>
<td>0.22</td>
<td>10.33</td>
<td>3.03</td>
<td>46.54</td>
<td>0.00</td>
<td>0.00</td>
<td>0.00</td>
<td>32.63</td>
<td>0.62</td>
<td>0.16</td>
<td>0.25</td>
<td>2.11</td>
<td>2.50</td>
<td>0.23</td>
<td>0.00</td>
</tr>
<tr>
<td>2</td>
<td>0.47</td>
<td>0.25</td>
<td>11.62</td>
<td>3.26</td>
<td>50.35</td>
<td>0.14</td>
<td>0.00</td>
<td>0.00</td>
<td>25.29</td>
<td>0.35</td>
<td>0.18</td>
<td>0.15</td>
<td>3.63</td>
<td>2.05</td>
<td>0.16</td>
<td>0.15</td>
</tr>
<tr>
<td>3</td>
<td>0.20</td>
<td>0.15</td>
<td>9.76</td>
<td>4.60</td>
<td>48.38</td>
<td>0.00</td>
<td>0.00</td>
<td>0.00</td>
<td>29.69</td>
<td>0.56</td>
<td>0.10</td>
<td>0.28</td>
<td>1.92</td>
<td>2.49</td>
<td>0.29</td>
<td>0.00</td>
</tr>
<tr>
<td>4</td>
<td>0.10</td>
<td>0.13</td>
<td>10.10</td>
<td>4.51</td>
<td>45.63</td>
<td>0.00</td>
<td>0.00</td>
<td>0.00</td>
<td>32.79</td>
<td>0.69</td>
<td>0.17</td>
<td>0.18</td>
<td>1.85</td>
<td>2.48</td>
<td>0.19</td>
<td>0.00</td>
</tr>
</tbody>
</table>

bdl = below detection limit.
4. Discussion

The nature and origin of fibrous PGM found in the Loma Larga ore deposit are enigmatic. Radial-fibrous internal textures of secondary PGM derived from the in situ alteration of magmatic laurite at low temperatures were reported from ophiolite-related chromitite samples from Greece [26], and similar secondary grains were also described by [27] from Turkish ophiolites. The latter authors demonstrated that dark zones in BSE images contain significant contents of Si and Mg (up to 9 and 10 at %, respectively), which accounted for sub-microscopic particles of chlorite or serpentine. Such secondary PGM grains with intimate intergrowths of Mg-silicates, believed to form during stages of serpentinization, occur in the study area ([8] and reference therein).

Recent studies have suggested that these Mg-silicates are not stable at higher levels of the Ni-laterite profile (close to the Mg-discontinuity) and are replaced by secondary Fe oxide(s) [20]. The observation of subsequent Fe-incorporation into secondary Ru-Os-Ir alloys via abiotic and/or biological processes has been used as an evidence for a supergene origin of ruthenian hexaferrum found in Ni-laterites from the Dominican Republic [20]. Here, the presented XRD data of one characteristic fibrous PGM with an isometric morphology suggest that such a PGM transformation from primary laurite toward ruthenian hexaferrum has occurred (Figure 3A; Tables 1 and 2). However, PGE-bearing nanofibers have not been observed in ruthenian hexaferrum grains from the study area before.

Complex textures such as PGM nanoparticles within film material (Figure 3I,J), as well as ragged morphologies of PGM nanoparticles resembling sponge spicules and an Ir-rich accumulation with a rather irregular shape (Figure 3K), point to continuous in situ neoformation on the surface of ruthenian
hexaferrum grains (Figure 3). Unfortunately, the nature of PGE-bearing nanofibers could not be determined in the present study and more analytical work is necessary to understand how associated PGM nanoparticles are formed. However, considering a starting material of secondary Ru-Os-Ir alloy hosting sub-microscopic fibrous serpentine minerals, it can be speculated that (bio-)weathering of these incorporated Mg-silicate remnants drives (i) the transformation of fibrous Mg-silicates into PGE-bearing nanofibers and (ii) the formation of PGM nanoparticles of different morphologies on the surface of the grain (Figure 3).

5. Conclusions

The PGM-rich “floating chromitite” found within the highest levels of the Loma Larga ore deposit is a suitable natural laboratory which can provide important new insights helping (i) to explain PGE mobilization in the surface environment; (ii) to understand transformation processes of hypogene PGM related to supergene (bio-)weathering and (iii) to find convincing evidence for supergene PGM neoformation and processes involved.

Acknowledgments: This research has been financially supported by FEDER Funds, the Spanish project CGL2015-65824-P and the Catalan project 2014-SGR-1661, as well as by a PhD grant to TA sponsored by the Ministerio de Economía y Competitividad (Spain). The authors gratefully acknowledge the help and hospitality extended by the staff of Falcondo mine (Falcondo Glencore). Excellent technical support during EPMA sessions by Xavier Llovet and during FE-SEM sessions by Eva Prats at the Serveis Científics i Tecnòlogics (University of Barcelona) is highly appreciated. Louis Cabri and Vladimir Rudashevsky are greatly thanked for their help during installation of the HS-11 laboratory in Barcelona. Special thanks are due to the Guest Editor of the Special Issue “Mineral Deposit Genesis and Exploration” Maria Economou-Eliopoulos. Three anonymous reviewers are thanked for their constructive criticism and valuable comments that have improved the manuscript.

Author Contributions: Francisco Longo contributed materials and logistics during field work; Thomas Aiglsperger, Joaquín A. Proenza, Mercè Font-Bardia, Salvador Galí, Josep Roqué, and Sandra Baurier-Aymat performed the experiments and analyzed the data; Thomas Aiglsperger and Joaquín A. Proenza wrote the paper.

Conflicts of Interest: The authors declare no conflict of interest.

References


© 2016 by the authors; licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC-BY) license (http://creativecommons.org/licenses/by/4.0/).